Scale Dependence of Rupture Barriers

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## Systematics of Large Transform Fault Earthquakes <u>-The G-R distribution is well predicted by simple thermal models</u>



Boettcher and Jordan, 2004 Bird et al., 2002

## Systematics of Large Transform Fault Earthquakes



$$N(M) = N_0 \left(\frac{M_0}{M}\right)^\beta \exp\left(\frac{M_0 - M}{M_c}\right)$$

$$M_c = CA_T^{3/4}; A_T = C_T L^{3/2} V^{-1/2}$$

Estimate of C from 1964-2001 predicts 2002-2009 data extremely well; Boettcher and McGuire, 2009;

Large (Mc) earthquakes do not rupture the entire fault.



The creeping area on the large OTFs put together is the equivalent of having 10 completely aseismic San Andreas Faults. Or roughly 100 'creeping sections'



### **Rupture Barriers and Seismic Coupling**



Stress is always high in the rupture barrier and earthquakes are always followed by post-seismic creep in the barrier region

Kaneko et al, 2010



# **Oceanic Transform Fault Friction**



Boettcher et al., 2007

Gabbro and Peridotite expected to be velocity weakening up to 500-600 °C ->matches depth extent for both interand intraplate earthquakes

#### Kohli et al., 2011

Serpentine expected to be velocity strengthening at low slip-rates but possibly undergo large amounts of dynamic weakening

#### Discovery and Gofar Transform Seismic Cycles

Mw 6 Ruptures repeat every ~5-6 years Indistinguishable centroids but seismic moments can vary by a factor of 2-3 between cycles. Overall, they are suggestive of fully coupled patches separated by creeping segments.

ClippertonMw ~6.5~19±3 yearsBlancoMw ~6.3~14±2 yearsGofarMw ~6.1~6±1 yearsDiscoveryMw ~6.0~5.5±1 yearsSiqueirosMw ~5.8~5.5±1 years



# The QDG Experiment

The first local observations of the end of a oceanic transform seismic cycle Quebrada, Discovery, Gofar Fault System



#### The Gofar Fault



The western asperity had M ~6.0 earthquakes in 1992, 1997, 2003, and 2008 and 2012. The eastern asperity had M ~6.0 earthquakes in 1997, 2002, and 2007 and maybe 2011.

 $\rightarrow$  The Foreshock zone is a long-lived rupture barrier.

350 300 250 Day of 2008 200 150 100 50 n -106.2 -105.8-106 Longitude

The failure of the rupture barrier may modulate the end of the seismic cycles.



## Physical Properties of the rupture barrier



Roland et al., 2012



M6 Rupture Zone





Van Avendonk et al. 1998

Full Couplingv M 6.6 to 6.8 earthquakes in 1911, 1936, 1957, 1978, and 1995



#### Cause of the Gofar Fault-Zone P-wave Velocity Reduction



Serpentine could explain the anomaly, but is not supported by the gravity field and Is unlikely to change on weekly time scales.

Roland et al., 2011

## Fault-Zone Velocity Changes During the Foreshock Sequence



The coda following S-waves that propagate vertically in the fault-zone (~4-12 Hz) shows clear stretching during the week long foreshock sequence.

### Fault-Zone Velocity Changes During the Foreshock Sequence



$$\frac{dt}{t} = -\frac{dv}{v}$$

From Pierre Gouédard

#### Space-time history of S-wave velocity changes



zone



Larger coseismic velocity drop in the mainshock zone than the rupture

Rapid temporal variations in the barrier during the foreshock sequence.

## Fault-Zone Velocity Changes During the Foreshock Sequence



The seismicity-rate jumps by about an order of magnitude during the foreshock sequence.

This would correspond to at least an order of magnitude jump in stressing-rate in the Dieterich 1994 rate state seismicity model.

## Summary

Rupture barriers to large earthquakes predominate on OTFs but are more complicated than a simple velocity strengthening zone

Are long-lived through multiple seismic cycles
Are able to nucleate intense micro-seismicity

> velocity weakening at least at many small asperities
•Show significantly more damage in the seismogenic zone than regions that host large earthquakes

> porosity of a few percent in the seismogenic zone
 Appear to have experienced a net drop in stress during the large rupture not the increase expected from simple models
 =>possible dynamic weakening?

•Show large time-dependence to their shear-wave velocities

=> high porosity /stress-dependence of velocity

=> coseismic damage may indicate damage

May also show lower apparent stress than mainshock zones (see Pamela Moyer's poster)

### What are the material properties of the rupture barrier/foreshock



## . Additional Insight from Gravity Data



Simple model of gravity signature over fault zone containing partially serpentinized peridotite



Free-air Anomaly does not show evidence for reduced density associated with altered mantle rocks A 0.5-5% porosity increase comprised of high aspect ratio cracks extending ~5-6 km bsf could explain the Vp anomaly.

What do the "mainshock" parts of the fault look like?



Active 0.85 Ma 1.3Ma



Roland et al., 2011

# **Velocity Changes**

#### **Foreshock Zone**

#### Mainshock Zone



From Pierre Gouédard

The key question is why does the foreshock zone show both an interseismic (positive) trend in velocity and multiple (negative) velocity decreases while the mainshock zone shows nothing?

# **Velocity Changes**

#### **Foreshock Zone**



From Pierre Gouédard

Velocity drops are correlated with earthquake swarms in multiple cases.