Frictional properties of Zuccale Fault rocks from room temperature to in-situ conditions:

Result from high strain rotary shear experiments

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Background – The Zuccale Fault



- Low angle normal fault, dip of $\sim 15^{\circ}$
- Accomodated ~6-8 km displacement
- Exhumed from max. 8 km depth
- Complex geometry & deformation history
- Production of talc from dolomite in high strain domains





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Why the Zuccale Fault?

-Analogue for currently active Alto-Tiburina Fault

- A) A constant seismicity rate 3.5 event/d, M_L<2.3, recorded during 8 month campaign (2000-2001).
- B) Composite focal mechanisms with a gently E-dipping plane.
- C) Multiple events with highly correlated waveforms.

Chiaraluce et al. JGR, 2007





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Zuccale fault_

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Problem & current aims

- Fault at high angle to σ_1 (vertical) => fault is weak !
- Weakness result of the presence of a phyllosilicate foliation
- But also (micro)seismicity !





Ikari et al, Geology, 2011

Previous work at T_{room} & dry => hydrothermal conditons:
-Weakening due to talc formation ?
-Effect of temperature on slip stability ?
-Effect of fluid pressure/normal stress on slip stability ?







Experimental apparatus Hydrothermal rotary shear apparatus at HPT Laboratory





-Ring-shaped samples 22/28 diameter -Internally heated pressure vessel -T_{max}=700 °C

- σ_{nmax} =~300 MPa, P_{fmax} =300 MPa -Pressure-compensated piston -Displacement rates 3 nm-300 µm/s -Infinite displacement



Sample material & procedure

- Simulated analogue of 80/20 wt % dolomite/quartz
- 4 powdered samples from the foliated fault core
- 1 powdered sample from a non-foliated portion (ZF05)
- Phyllosilicate content variable from ~6 wt% up to 58 wt%

Sample	Ca <mark>lcite</mark>	Tremolite	Talc	Clays	Dolomite
ZF01	<mark>42</mark>	0	56	2	0
ZF02	<mark>23</mark>	42	22	13	0
ZF03	<mark>78</mark>	0	16	5	0
ZF04	4 <mark>3</mark>	46	0	6	0
ZF05	2 <mark>9</mark>	0	7	0	64

- ~ 1 mm starting thickness
- Run-in at 10 μ m/s, followed by velocity-steps from 0.3 100 μ m/s
- Velocity-dependence of friction expressed as (a-b) = Δμ_{ss}ln(v/v₀),
 (a-b) < 0 => potentially unstable (seismic) slip







Rate and state friction



Coupled with Dieterich evolution equation





Results: comparison with previous experiments, talc-rich sample ZF01



- Presence of fluids at pressure has only small effect on μ
- Velocity-strengthening enhanced







General decrease in friction with increasing phyllosilicate-content
Some weakening with temperature







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- (a-b) more positive for samples with more phyllosilicates
- Negative (a-b) at T=150 °C & 300°C and low v
- In some cases unstable slip observed
- (a-b) increases with increasing velocity





Results – effect of $P_f \& \sigma_n^{eff}$



• Only small effect on friction coefficient in all samples





Results – effect of $P_f \& \sigma_n^{eff}$

0.85

Sample ZF04 at T=300 °C



Fluid pressure or effective stress?



- Smaller (a-b) when $P_f \ge 120 \text{ MPa}$
- Str<mark>ain has little effect</mark>



























Significant change in (a-b) !







- 1. Large variation in strength of fault units as a function of the amount of reaction products (notably talc).
- 2. The velocity dependence of friction of Zuccale fault rocks under hydrothermal conditions depends strongly on the velocity.
- 3. Strong portions of the fault show the potential for unstable, seismic slip at elevated temperature and low sliding velocity.
- 4. Negative velocity-dependence of friction is more pronounced at elevated fluid pressures and low effective normal stresses.
- 5. Change in (a-b) due to mineral reactions at low velocity
- 6. Velocity dependence of friction controlled by a time-dependent, thermally activated, fluid-assisted process ? Similar velocity-dependent behaviour observed in calcite gouges at 100-150 °C (Verberne et al., BSSA, 2011) Stability controlled by carbonate-content ?







Thank you !





