UNIVERSITY of NEW HAMPSHIRE

Pamela Moyer¹, Margaret Boettcher¹, Jeffrey McGuire², and John Collins²

Motivation

On Gofar Transform Fault on the East Pacific Rise, we observe relatively stable rupture patches over four seismic cycles. These rupture patches are separated by M_w 6.0 earthquakes that occur approximately every 5 years. The rupture patches are separated by regions of very low coupling which host swarms of microearthquakes. In 2008, McGuire and coworkers captured the end of a seismic cycle with an ocean bottom seismic array. The dataset includes an extensive foreshock sequence (right, yellow dots), the mainshock and aftershocks (right, red dots), and an earthquake swarm that occurred three months after the mainshock (right, blue dots).

We investigate source parameters of small to intermediate magnitude earthquakes $(3.0 \le M_{w} \le 4.7)$ along the Gofar transform fault to assess differences in earthquake properties between rupture patches and rupture barriers.

Far right:

A. Locations of over 20,000 earthquakes that occurred along Gofar transform fault from August - December 2008. Ocean bottom seismometers are noted by triangles and stars. The epicentre of the largest (M 5.2) aftershock and the centroids of the 2008 M_w 6.0 and 2007 M_w 6.2 earthquakes are shown as brown, orange, and blue circles, respectively.

B. Earthquake temporal distribution showing the rate of seismicity in the foreshock, mainshock and aftershock, and swarm zones. **C.** Along strike variations in earthquake depth in the foreshock, mainshock and aftershock, and swarm zones [McGuire et al., 2012].

Map and space-time evolution of large earthquakes on the Quebrada, Discovery, and Gofar transform faults (after McGuire 2008). Earthquakes with $M_{\rm m} > 5.0$ since 1990 are shown as circles at their locations in the Gobal CMT catalog. Events with overlapping rupture areas (defined as relative centroid locations < 5 km, see McGuire 2008) are shown in a constant color. Vertical gray lines show the locations of segment boundaries defined by intra-transform spreading centers [McGuire e. al., 2012].



We calculate radiated seismic energy using an omega-squared source model, where the corner frequency is derived from an empirical Green's function (EGF) technique. We follow the method of Viegas et al., [2010] to find good EGF pairs and compare the results of stacking "high-quality" spectral ratios (as defined by the variance in the fit to the corner frequencies of both events) to stacking all station, component data.



Due to noise and resonance in the OBS data, most of our results come from one station. Often, only one component ratio fit to stacked data from one station passes our quality tests. The question of how to deal with limited spectral data arises. Is only one "high-quality" ratio enough for robust results? Should all data be stacked? What other quality tests should be performed?





McGuire, Jeffrey J., John A. Collins, Pierre Gouédard, Emily Roland, Dan Lizarralde, Margaret S. Boettcher, Mark D. Behn and Robert D. van der Hilst (2012), Variations in earthquake rupture properties along the Gofar transform fault, East Pacific Rise, Nature Geo. 5, 336-341, doi:10.1038/ngeo1454. McGuire, J. J. (2008), Seismic Cycles and Earthquake Predictability on East Pacific Rise Transform Faults, Bull. Seis. Soc. Am., 98, 1067-1084, doi: 1010.1785/0120070154.

Prieto, G.A., R. L. Parker, F. L. Vernon III (2009), A Fortran 90 library for multitaper spectrum analysis, Comput. Geosci., 35, 1701–1710, doi:10.1016/j.cageo. Viegas, Gisela, Rachel E. Abercrombie, and Won-Young Kim (2010), The 2002 M5 Au Sable Forks, NY, earthquake sequence: Source scaling relationships and energy budget, J. Geophys. Res., 115, B07310, doi:10.1029/2009JB006799.

Yang, Hongfeng, Lupei Zhu, and Risheng Chu (2009), Fault-Plane Determination of the 18 April 2008 Mount Carmel, Illinois, Earthquake by Detecting and Relocating Aftershocks, Bull. Seismol. Soc. Am., 99, 3413 – 3420, doi: 10.1785/0120090038