



# P-wave Attenuation with Implications for Earthquake Early Warning

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# Overview

- P-wave attenuation equations
  - Used for real-time magnitude estimations
  - Empirically derived (up to now)
- Deriving P-wave theoretical attenuation laws
- Relying on well established models
  - Brune omega-squared model
  - Parseval's theorem
- Verification against Japan and California data
- Implications for Earthquake Early Warning (EEW)
- Prospective

# Theory

## Source Spectra

$\omega = 2\pi f$ : Frequency  
 $\Omega(\omega)$ : Displacement spectral amplitude  
 $\dot{\Omega}(\omega)$ : Displacement spectral amplitude  
 $\Omega_0$ : Low Frequency plateau  
 $\omega_0 = 2\pi f_0$ : Corner frequency

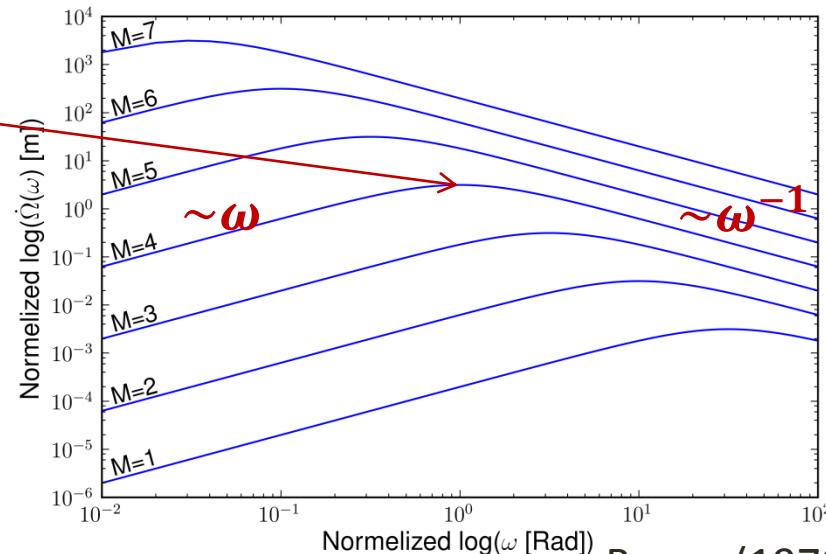
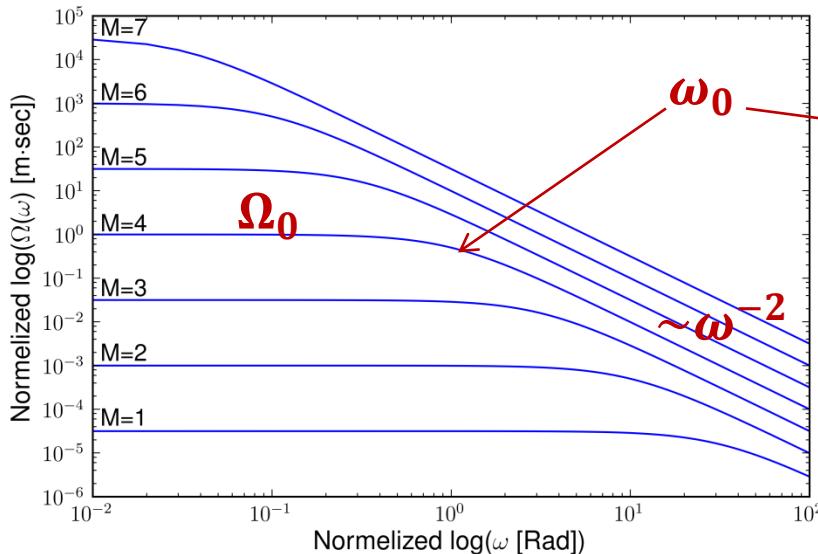
- Brune omega-squared model

- Displacement

$$\Omega(\omega) = \frac{\Omega_0}{1 + \left(\frac{\omega}{\omega_0}\right)^2}$$

- Velocity

$$\dot{\Omega}(\omega) = \frac{\omega \Omega_0}{1 + \left(\frac{\omega}{\omega_0}\right)^2}$$



Brune (1970)

# Theory

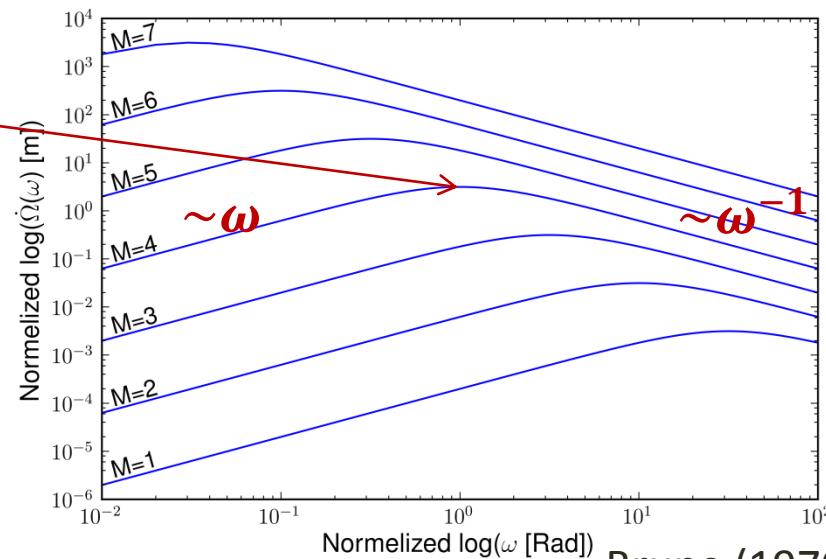
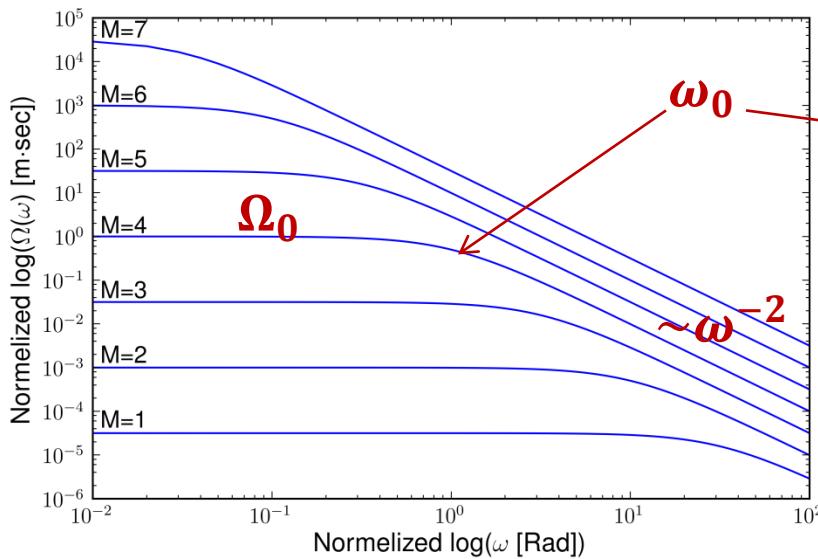
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Brune (1970)

# Theory

## Source Spectra

$M_0$ : Seismic moment  
 $U_{\varphi\theta}$ : Radiation pattern  
 $F_S$ : Free surface correction  
 $\rho$ : Density  
 $C_P$ : P-wave velocity  
 $R$ : Hypocentral distance

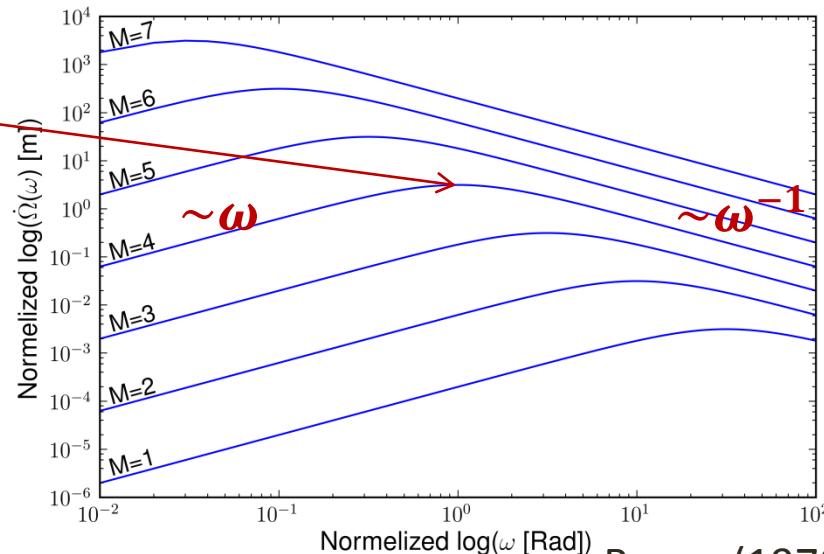
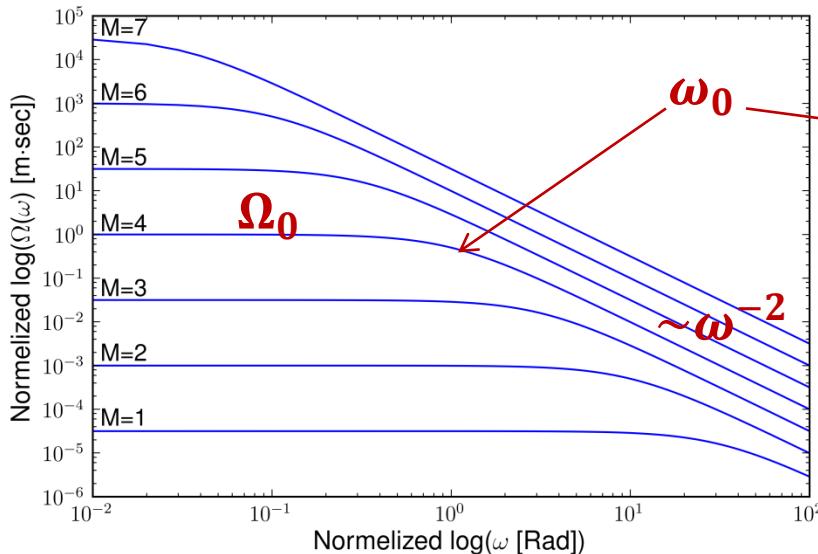
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$$\Omega(\omega) = \frac{\Omega_0}{1 + \left(\frac{\omega}{\omega_0}\right)^2}$$

$$\Omega_0 = \frac{M_0 U_{\varphi\theta} F_S}{4\pi\rho C_P^3 R}$$

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Brune (1970)

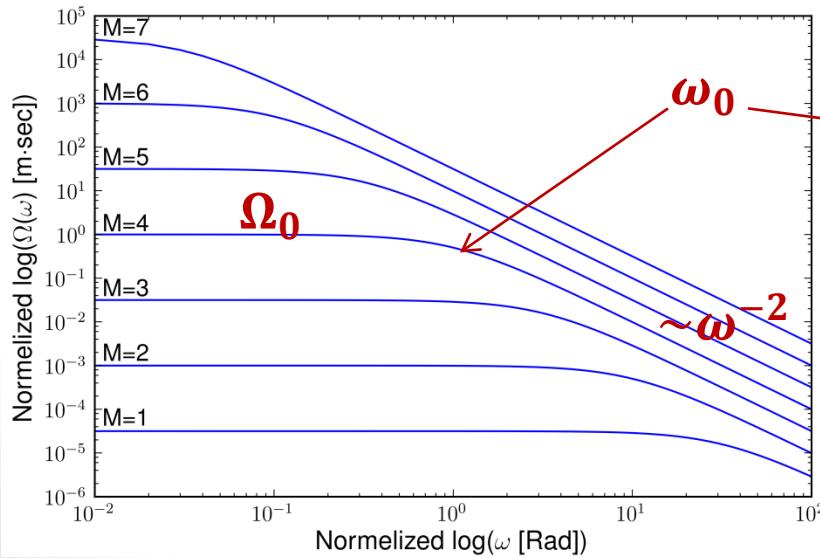
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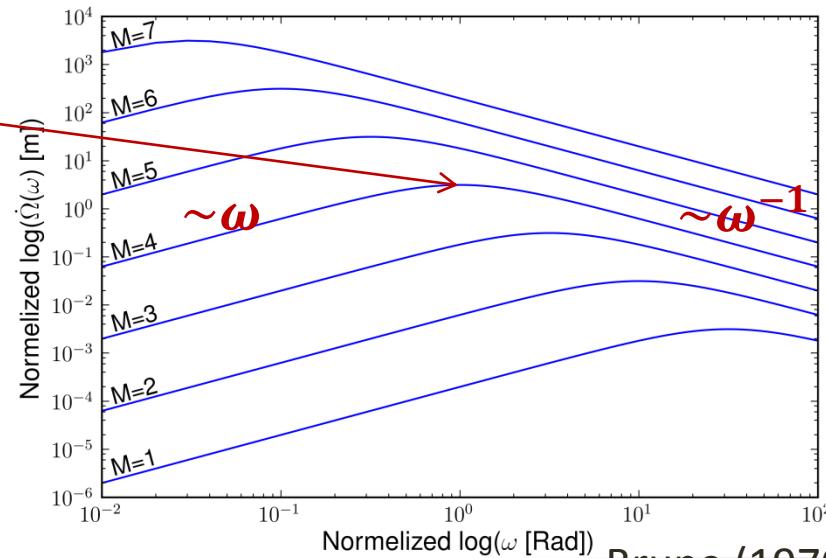
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- Velocity

$$\dot{\Omega}(\omega) = \frac{\omega\Omega_0}{1 + \left(\frac{\omega}{\omega_0}\right)^2}$$

$$f_0 = \frac{kC_S}{r}$$



$M_0$ :	Seismic moment
$U_{\varphi\theta}$ :	Radiation pattern
$F_s$ :	Free surface correction
$\rho$ :	Density
$C_P$ :	P-wave velocity
$R$ :	Hypocentral distance
$k$ :	Constant
$C_S$ :	S-wave velocity
$r$ :	Rupture radius

Brune (1970)

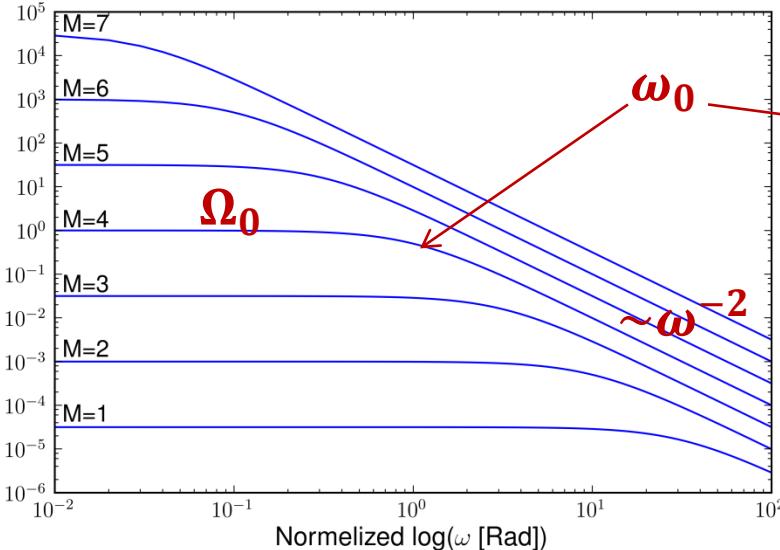
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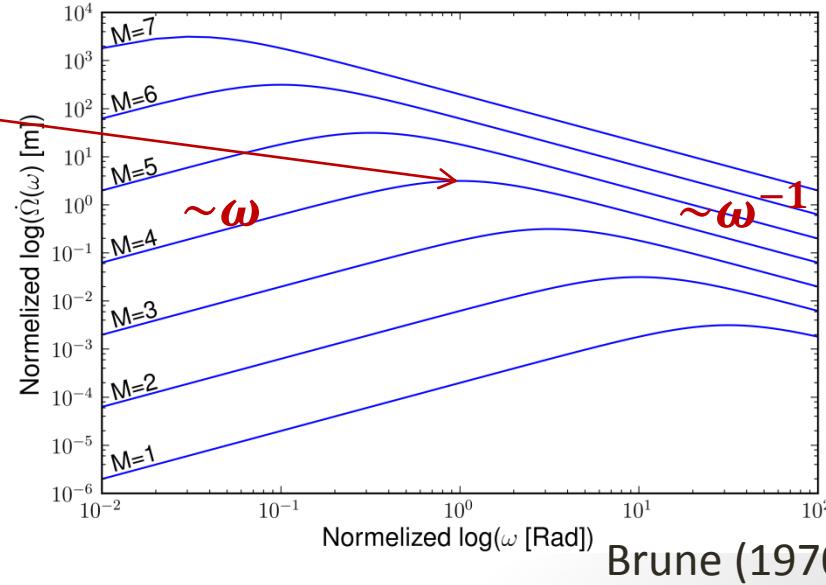
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$$\dot{\Omega}(\omega) = \frac{\omega\Omega_0}{1 + \left(\frac{\omega}{\omega_0}\right)^2}$$

$$f_0 = \frac{kC_S}{r}$$

$$M_0 = \frac{16}{7} \Delta\tau r^3$$

Eshelby (1957)



$M_0$ :	Seismic moment
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$F_s$ :	Free surface correction
$\rho$ :	Density
$C_P$ :	P-wave velocity
$R$ :	Hypocentral distance
$k$ :	Constant
$C_S$ :	S-wave velocity
$r$ :	Rupture radius
$\Delta\tau$ :	Stress drop

Brune (1970)

# Theory

## Frequency Domain Root-Mean-Square

- Time domain rms:

$$y_{rms} = \sqrt{\frac{1}{T_2 - T_1} \int_{T_1}^{T_2} |y(t)|^2 dt}$$

- Parseval's theorem:

$$\int_{-\infty}^{\infty} |y(t)|^2 dt = \int_{-\infty}^{\infty} |Y(f)|^2 df$$

# Theory

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- Time Window:

$$T_{S-P} = R \left( \frac{1}{C_S} - \frac{1}{C_P} \right) \stackrel{\text{def}}{=} R\eta$$

# Theory

## Attenuation Law Derivation

$$\Omega(\omega) = \frac{\Omega_0}{1 + \left(\frac{\omega}{\omega_0}\right)^2} \quad \dot{\Omega}(\omega) = \frac{\omega\Omega_0}{1 + \left(\frac{\omega}{\omega_0}\right)^2}$$

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$$f_0 = \frac{kC_s}{r}$$

$$\Omega_0 = \frac{M_0 U_{\varphi\theta} F_s}{4\pi\rho C_p^3 R}$$

$$M_0 = \frac{16}{7} \Delta\tau r^3$$

$$T_{S-P} = R\eta$$

# Theory

## Attenuation Law Derivation

$$d_{rms} = \Omega_0 \sqrt{\frac{\pi}{2} \frac{f_0}{T_{S-P}}}$$

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$$d_{rms} = \epsilon \left( \frac{16}{7} \Delta \tau \right)^{1/6} M_0^{5/6} \frac{1}{R^{3/2}}$$

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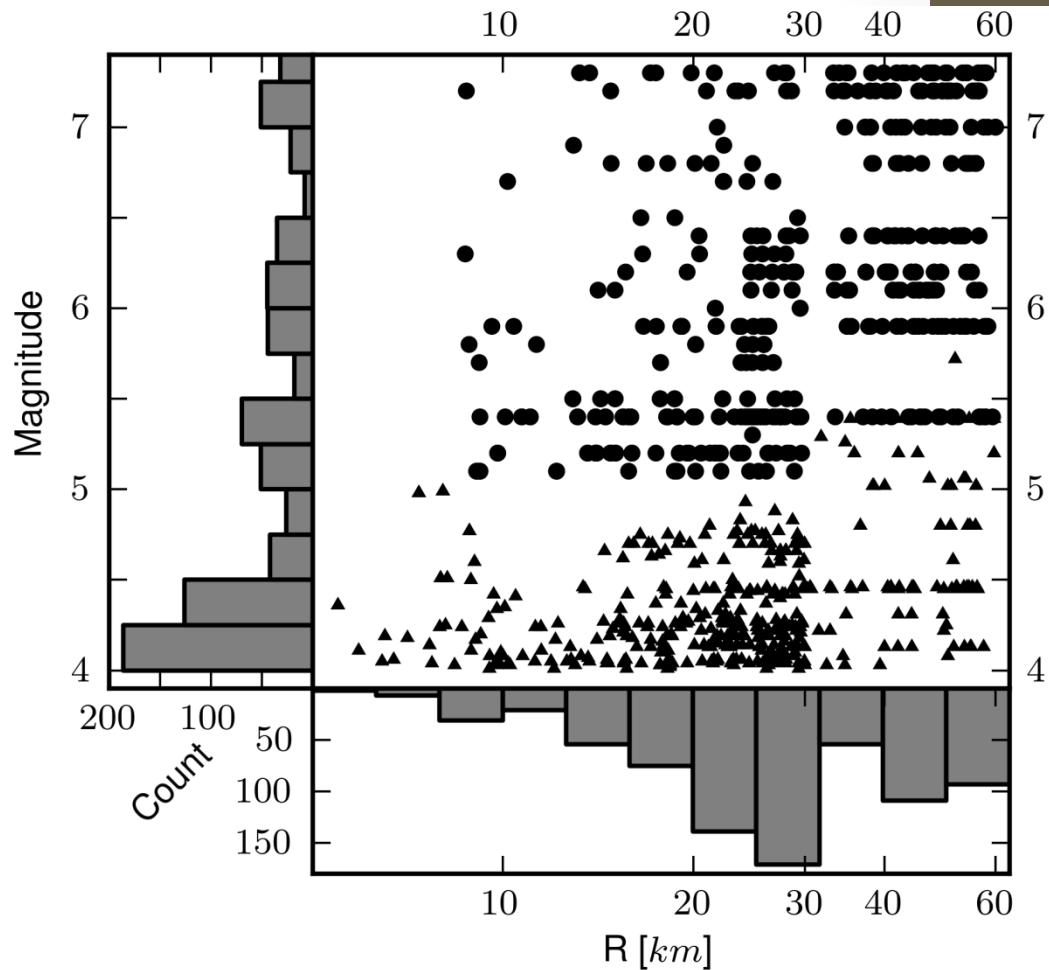
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$$T_{S-P} = R\eta$$

# Validation

## Data

- Japan
  - K-net + KiK surface
  - 353 accelerograms
  - 42 earthquakes
  - $5.1 < M < 7.3$
- California
  - CISN
  - 403 velocity seismograms
  - 120 earthquakes
  - $4 < M < 5.7$



# Validation

## Parameters

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$$\epsilon = \frac{U_{\varphi\theta} F_s}{4\pi\rho C_P^3} \sqrt{\frac{\pi k C_S}{2\eta}}$$

- $U_{\varphi\theta} = 0.52$
- $F_s = 2$
- $\rho = 2600 \text{ kg/m}^3$
- $C_S = 3200 \text{ m/s}$
- $C_P = 5333 \text{ m/s}$
- $k = 0.32$
- $\eta = 1/8 \text{ s/km}$
- $\epsilon = 753 \cdot 10^{-15} \text{ 1/Pa}$

# Validation

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- $\epsilon = 753 \cdot 10^{-15} \text{ 1/Pa}$
- $M_o, R - \text{Catalog}$
- $rms = \sqrt{\sum_{i=1}^h [AZ_i^2 + AE_i^2 + AN_i^2]/h}$

# Validation

## Stress Drop Determination

$$d_{rms} = \epsilon \left( \frac{16}{7} \Delta \tau \right)^{1/6} M_0^{5/6} \frac{1}{R^{3/2}}$$

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$$\Delta \tau = \frac{7}{16} M_0 \left( \frac{1}{k C_S} \cdot \frac{1}{2\pi} \frac{v_{rms}}{d_{rms}} \right)^3$$

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- Recalling Eshelby:

$$M_0 = \frac{16}{7} \Delta \tau r^3$$

$$1/r$$

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- And brune:

$$f_0 = \frac{k C_S}{r}$$

$$f_0$$

# Validation

## Stress Drop Determination

- Condition:

$$T_{S-P} > T_r$$

$$T_r = \frac{2r}{0.9C_S}$$

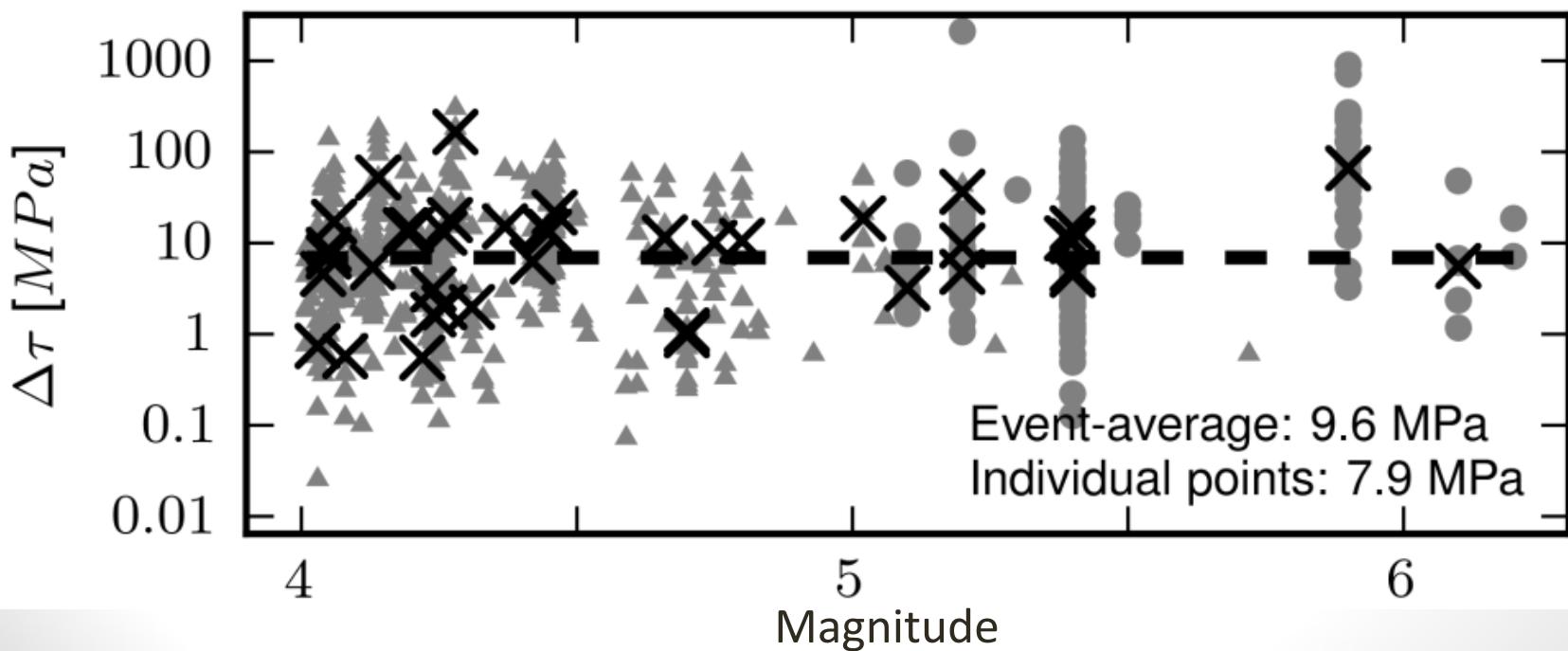
- 67% of data satisfies the condition

# Validation

## Stress Drop Determination

$$\Delta\tau = \frac{7}{16} M_0 \left( \frac{1}{kC_S} \cdot \frac{1}{2\pi} \frac{v_{rms}}{d_{rms}} \right)^3$$

Median:  $\Delta\tau = 7.9 \text{ MPa}$

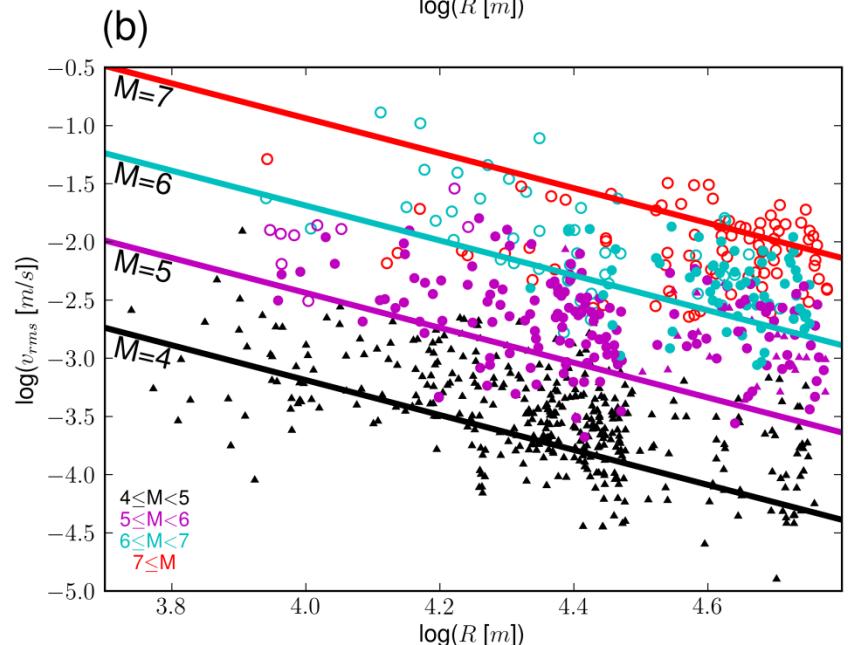
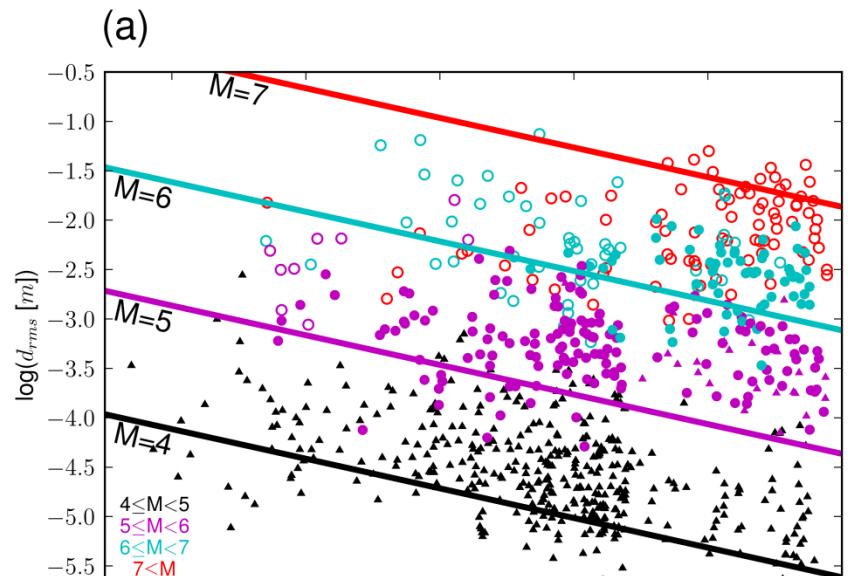


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## Observed Attenuation

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- Displacement

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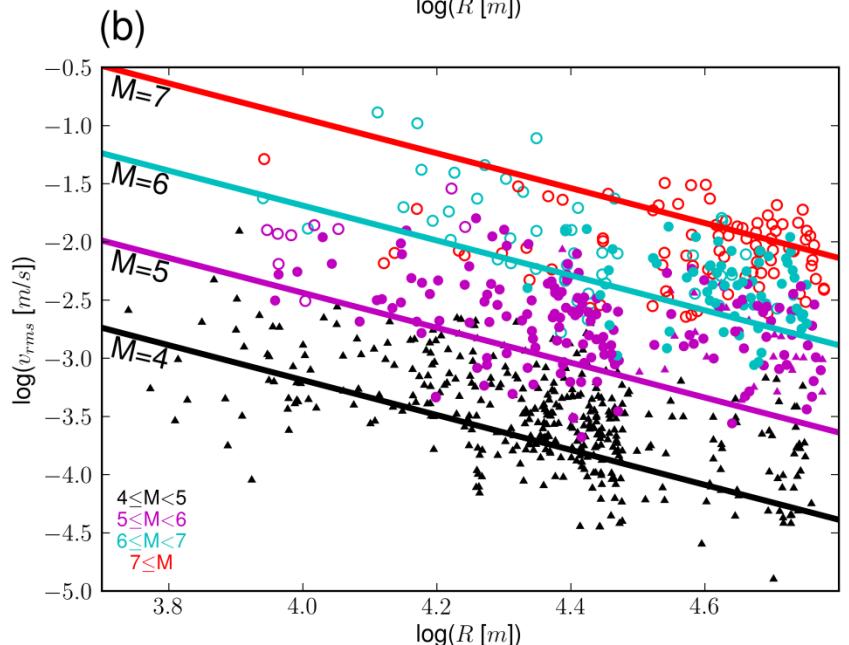
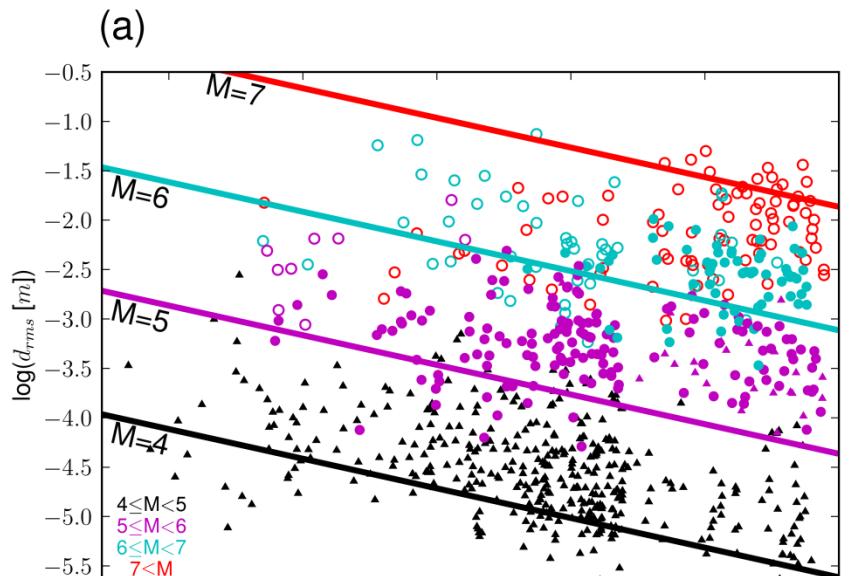
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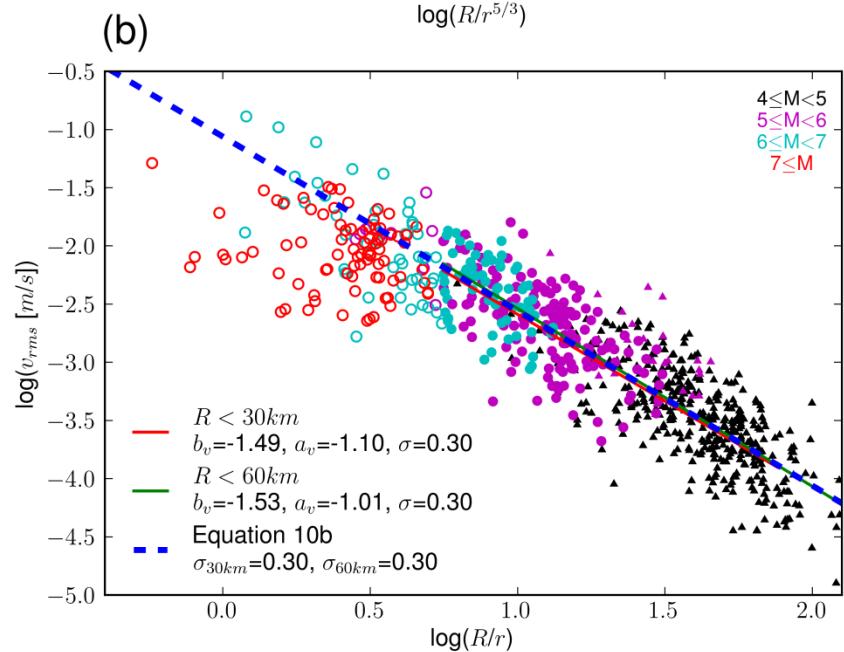
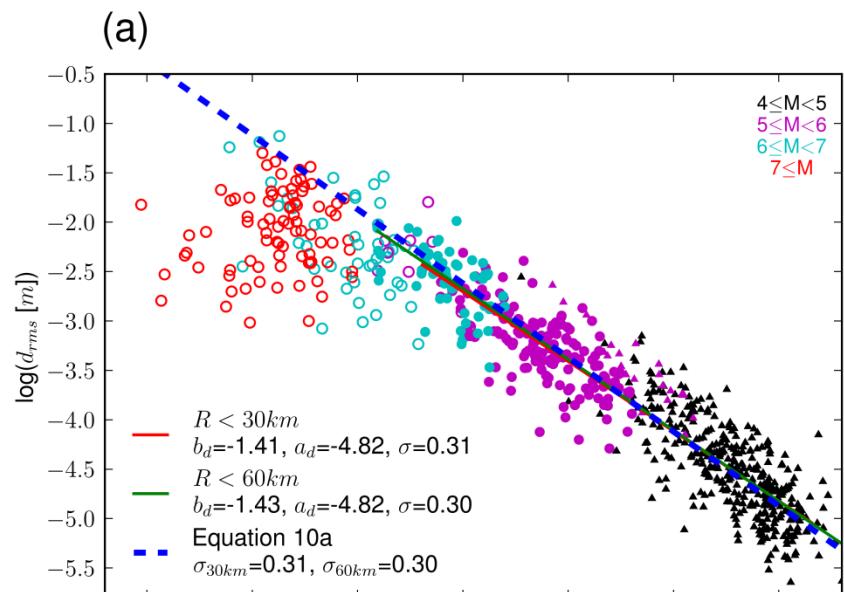


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$$d_{rms} = \epsilon \frac{16}{7} \Delta\tau \left( \frac{r^{5/3}}{R} \right)^{3/2}$$

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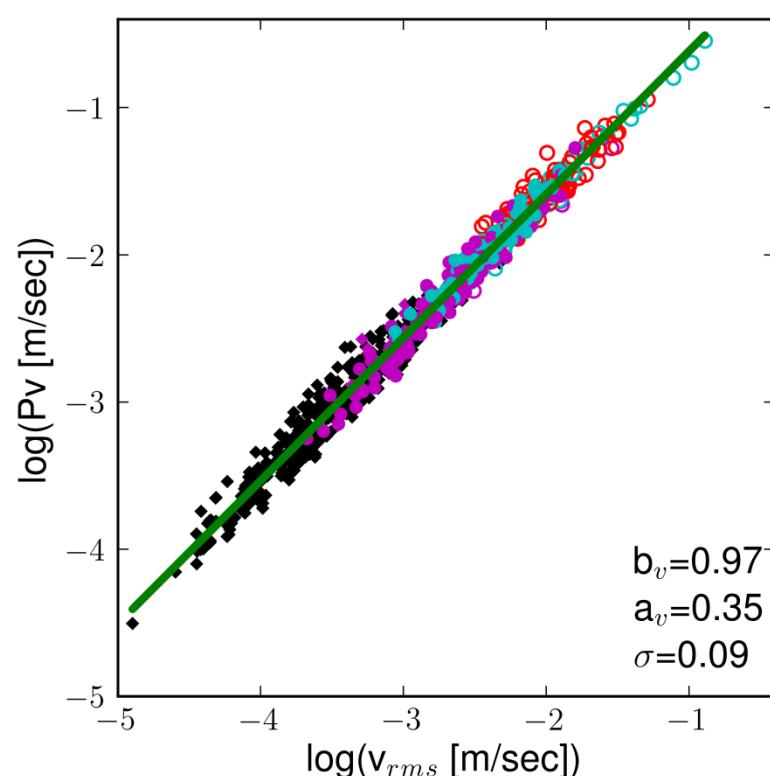
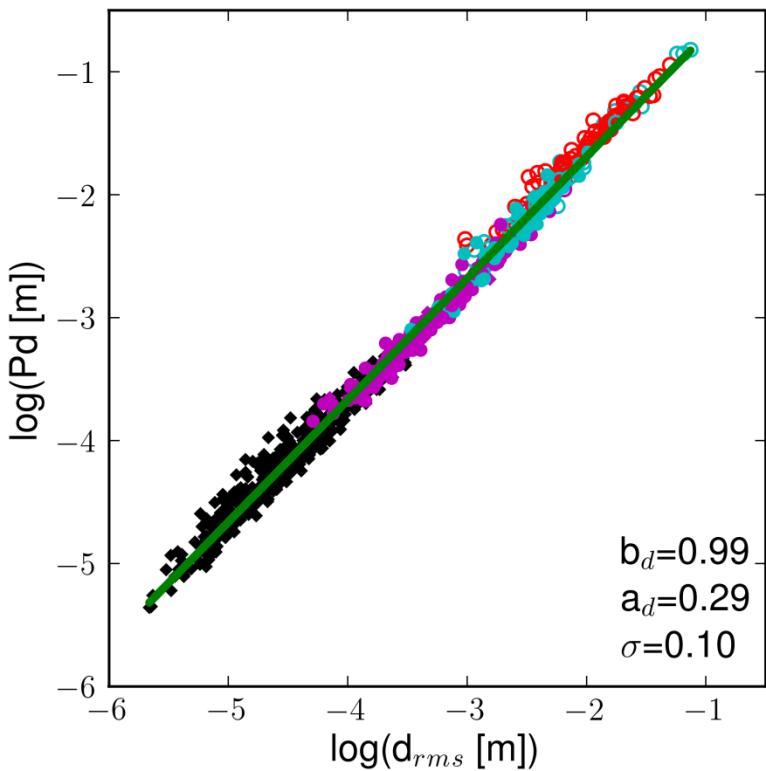


# Validation

## Implications for EEW: *rms* to Peak Amplitude

- Displacement
- $Pd = 2.0(\pm 0.5)d_{rms}$
- Velocity
- $Pv = 2.3(\pm 0.5)v_{rms}$

$$Pd \text{ or } Pv = \max \left[ \sqrt{AZ_i^2 + AE_i^2 + AN_i^2} \right]$$



# Validation

## Implications for EEW: $\tau_c$

- Characteristic period
- Real-time magnitude proxy
  - $\tau_c - Pd$  EEWs

$$\tau_c \stackrel{\text{def}}{=} 2\pi \sqrt{\frac{\int_{t_0}^{t_0+\Delta t} u^2 dt}{\int_{t_0}^{t_0+\Delta t} v^2 dt}} = 2\pi \frac{d_{rms}}{v_{rms}}$$

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$$f_0 = \frac{1}{2\pi} \frac{v_{rms}}{d_{rms}}$$

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$$\Delta\tau = \frac{7}{16} M_0 \left( \frac{1}{kC_S} \cdot \frac{1}{2\pi} \frac{v_{rms}}{d_{rms}} \right)^3$$

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- Real-time magnitude proxy

- $\tau_c$  – Pd EEWs

$$\tau_c \stackrel{\text{def}}{=} 2\pi \sqrt{\frac{\int_{t_0}^{t_0+\Delta t} u^2 dt}{\int_{t_0}^{t_0+\Delta t} v^2 dt}} = 2\pi \frac{d_{rms}}{v_{rms}}$$

$$\tau_c = \frac{1}{f_0}$$

$$f_0 = \frac{1}{2\pi} \frac{v_{rms}}{d_{rms}}$$

$$\tau_c = \frac{1}{kC_s} \sqrt[3]{\frac{7}{16} \frac{M_0}{\Delta\tau}}$$

$$\Delta\tau = \frac{7}{16} M_0 \left( \frac{1}{kC_s} \cdot \frac{1}{2\pi} \frac{v_{rms}}{d_{rms}} \right)^3$$

# Discussion

## Magnitude Estimations

- Attenuation laws show:

$$d_{rms} \propto \Delta\tau^{1/6} M_0^{5/6} R^{-3/2}$$

$$v_{rms} \propto \Delta\tau^{1/2} M_0^{1/2} R^{-3/2}$$

# Discussion

## Magnitude Estimations

$$M_0 \propto d_{rms}^{1.2} \Delta\tau^{-0.2} R^{1.8}$$

$$M_0 \propto v_{rms}^2 \Delta\tau^{-1} R^3$$

# Discussion

## Magnitude Estimations

$$M_0 \propto d_{rms}^{1.2} \Delta\tau^{-0.2} R^{1.8}$$

$$M_0 \propto v_{rms}^2 \Delta\tau^{-1} R^3$$

$$\tau_c \longleftrightarrow M_0 \propto v_{rms}^3 d_{rms}^{-3} \Delta\tau$$

# Discussion

## Magnitude Estimations

$$M_0 \propto d_{rms}^{1.2} \Delta\tau^{-0.2} R^{1.8}$$

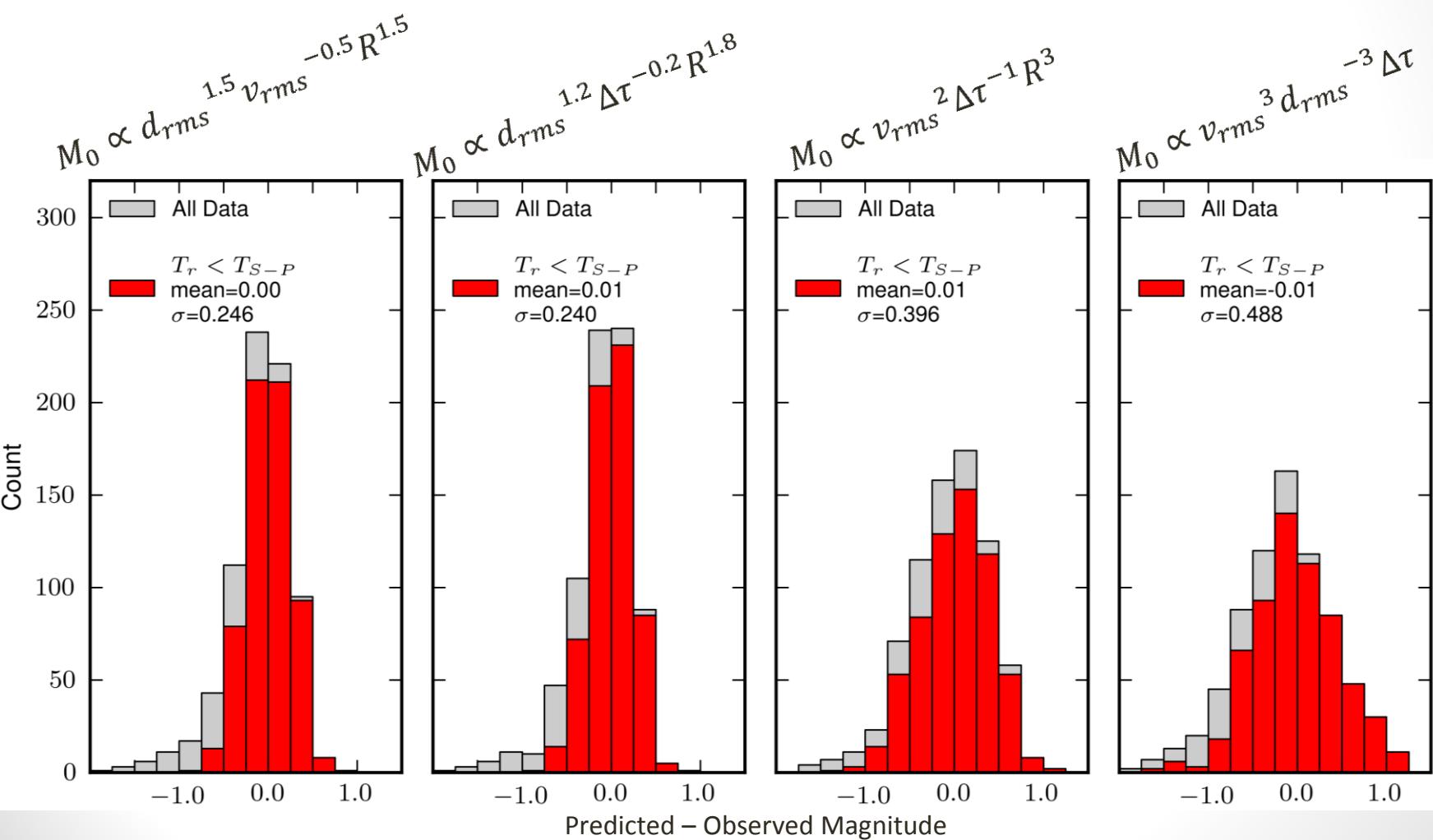
$$M_0 \propto v_{rms}^2 \Delta\tau^{-1} R^3$$

$$\tau_c \longleftrightarrow M_0 \propto v_{rms}^3 d_{rms}^{-3} \Delta\tau$$

$$M_0 \propto d_{rms}^{1.5} v_{rms}^{-0.5} R^{1.5}$$

# Discussion

## Magnitude Estimations



# Conclusion

- New **theoretical** attenuation laws
  - Provide source parameter weightings
  - Better stress drop and seismic moment understanding
- New stress drop estimation schemes
- Verified for Japan and California
- Earthquake Early Warning (EEW) implications
  - Little calibration needed
  - $rms$  – peak amplitude relation
  - $d_{rms}^{1.5} / v_{rms}^{0.5}$  magnitude estimation
  - $\tau_c = 1/f_0$
- prospective
  - Stress drop estimation
  - Ground motion prediction equation